



Government of **Western Australia**
Department of **Mines and Petroleum**

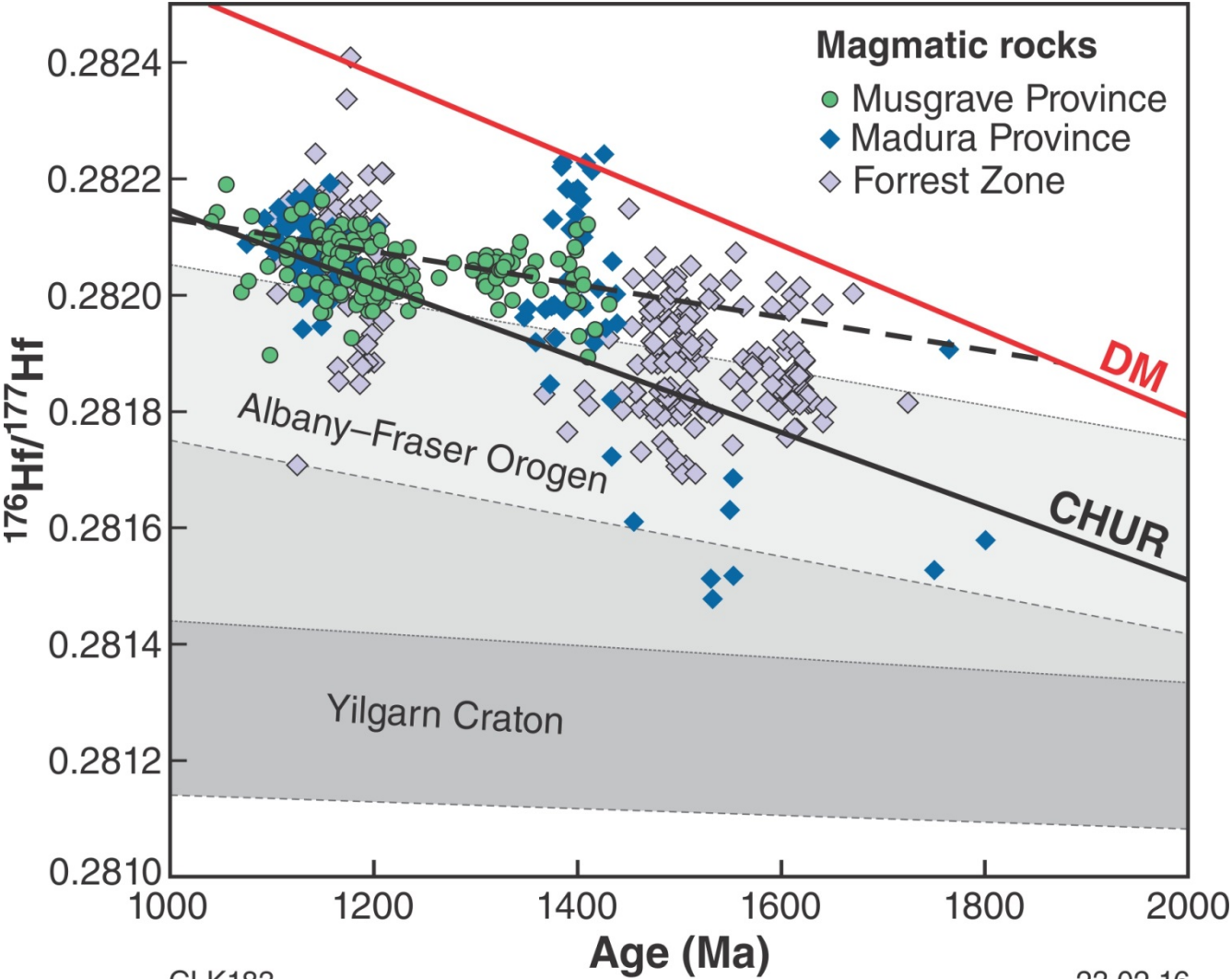
Exposing the Eucla basement: what separates the Albany–Fraser Orogen and the Gawler Craton?

PART 2

RH Smithies, CV Spaggiari, CL Kirkland, MTD Wingate
and RN England

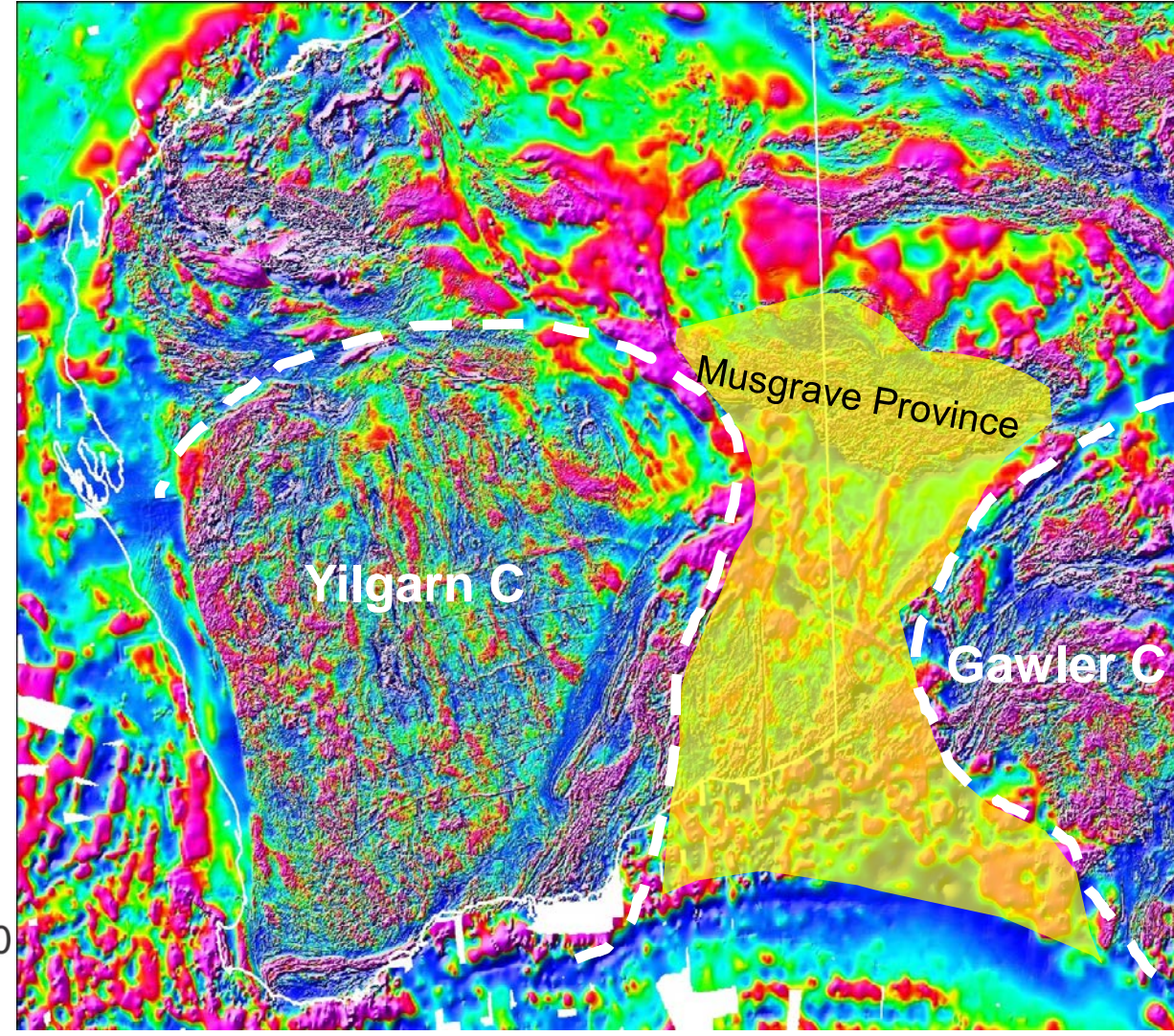


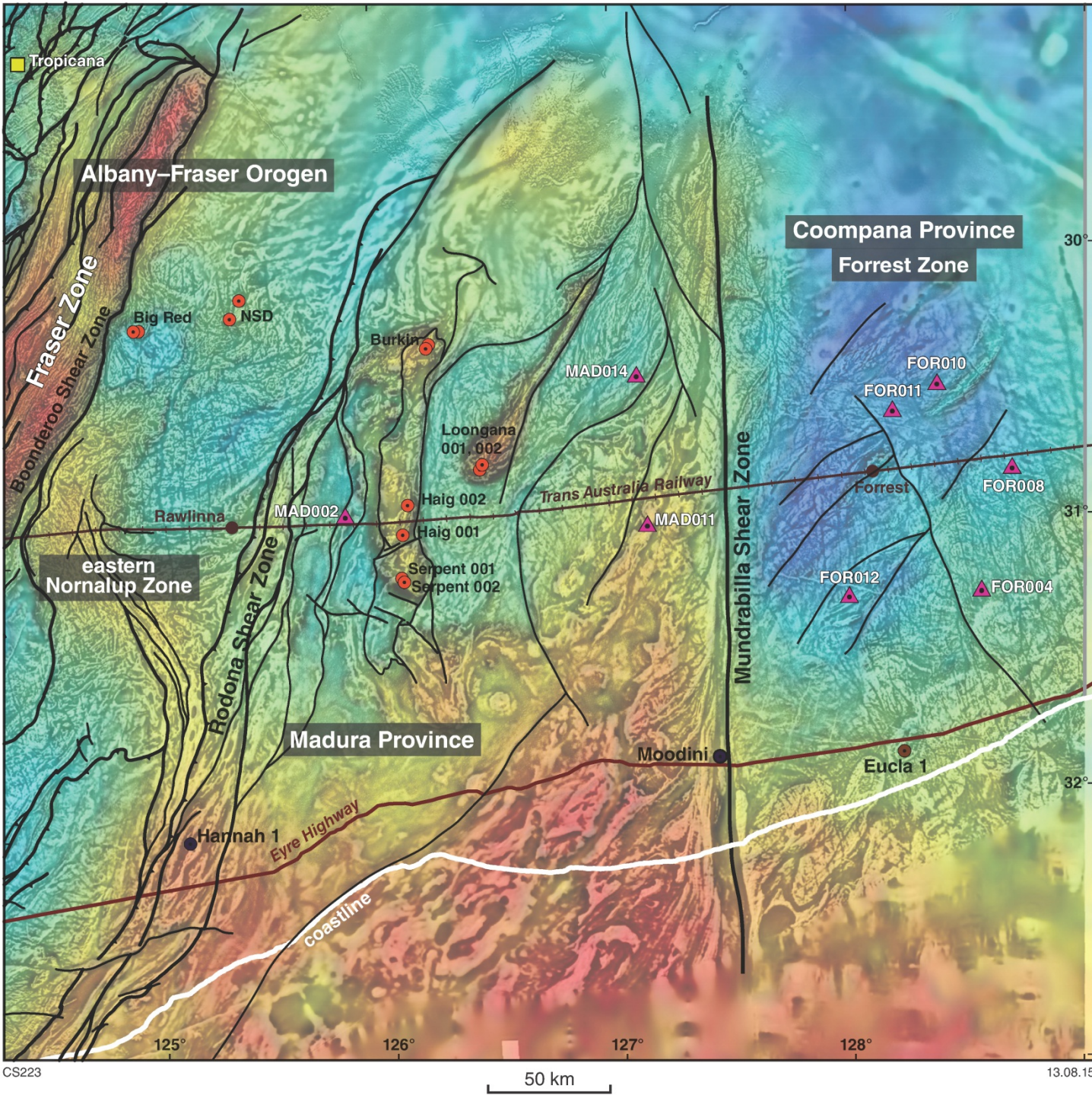
Forrest Zone, Madura Province and Musgrave Province developed on the same piece of juvenile mafic (oceanic!) crust created at c. 1950 Ma



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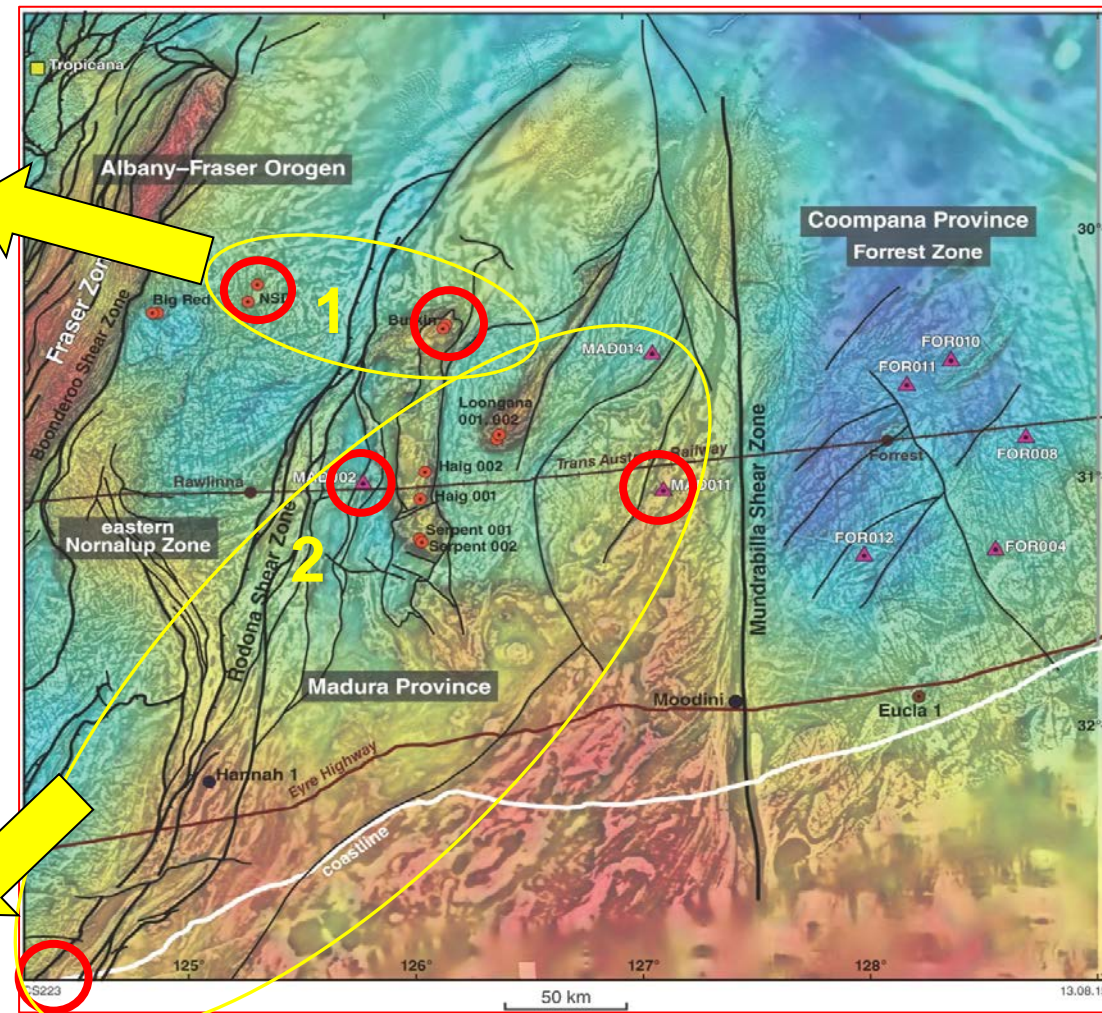
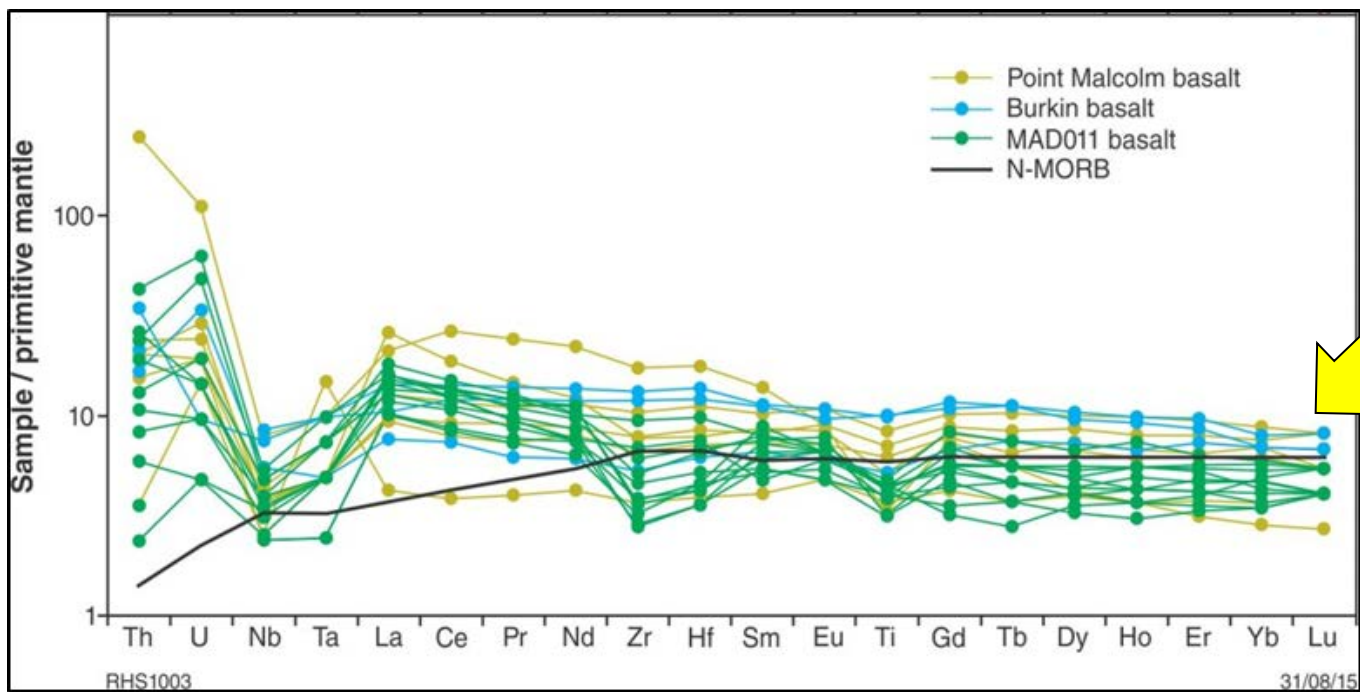
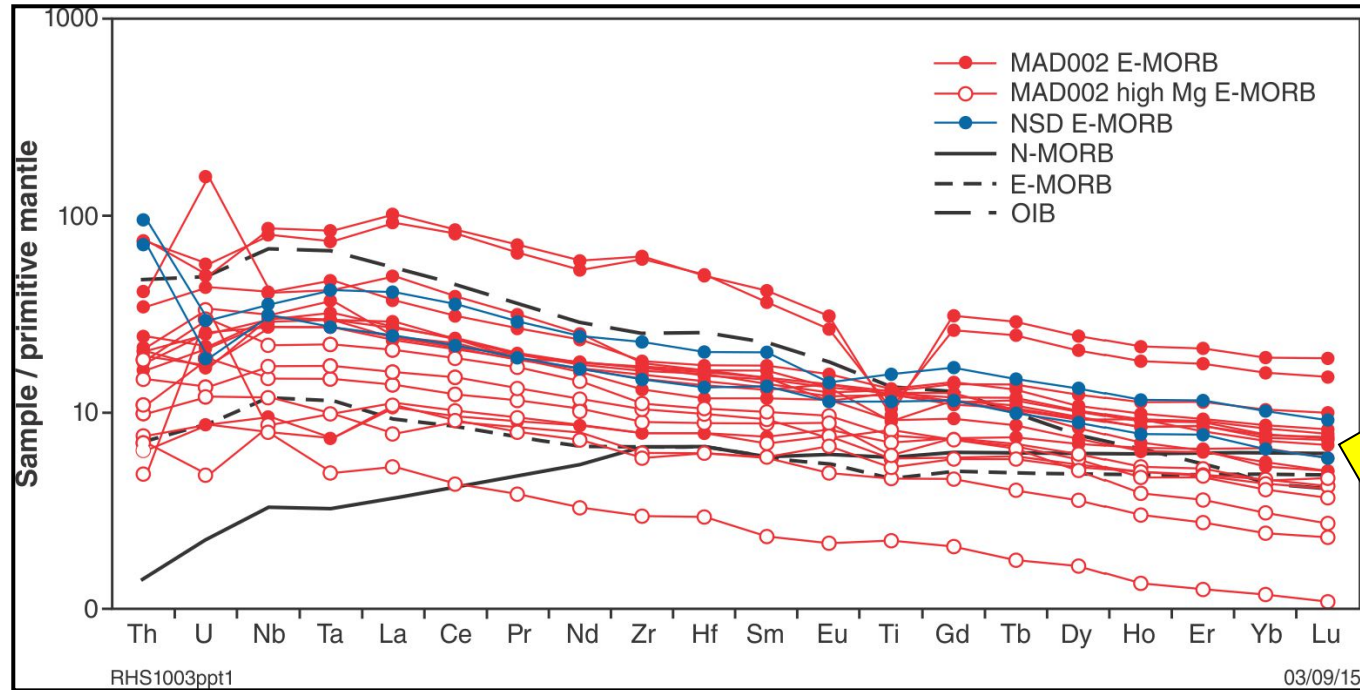


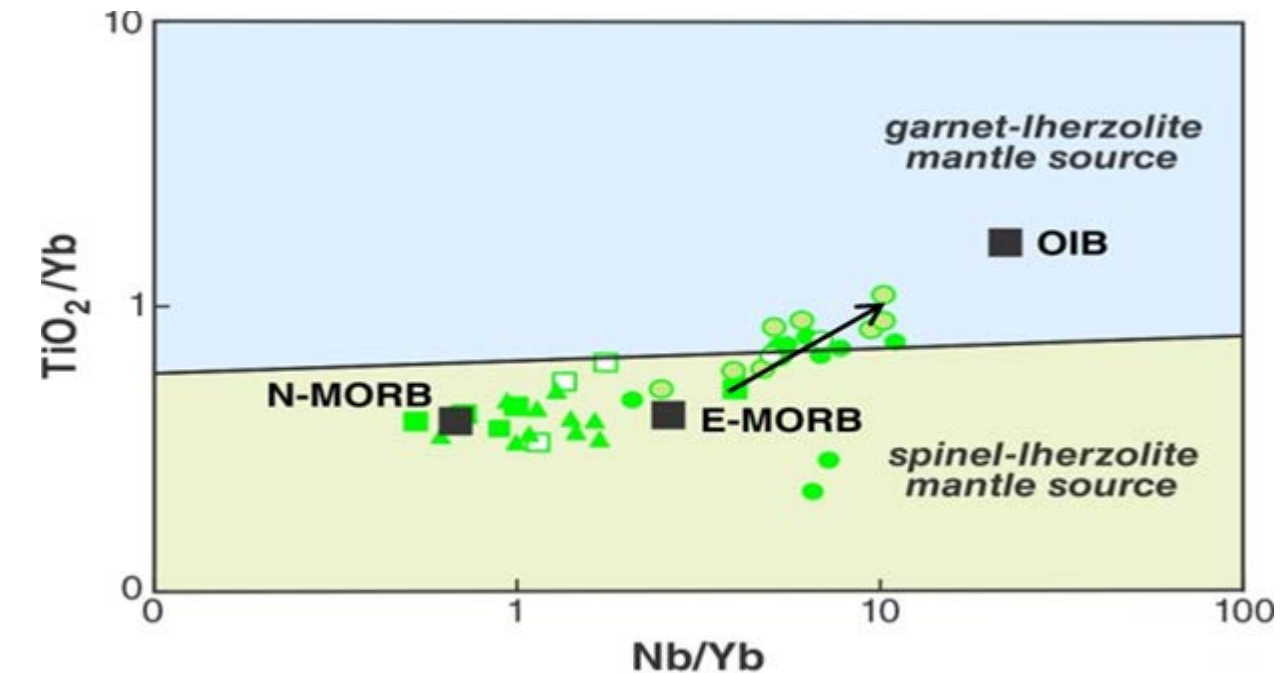
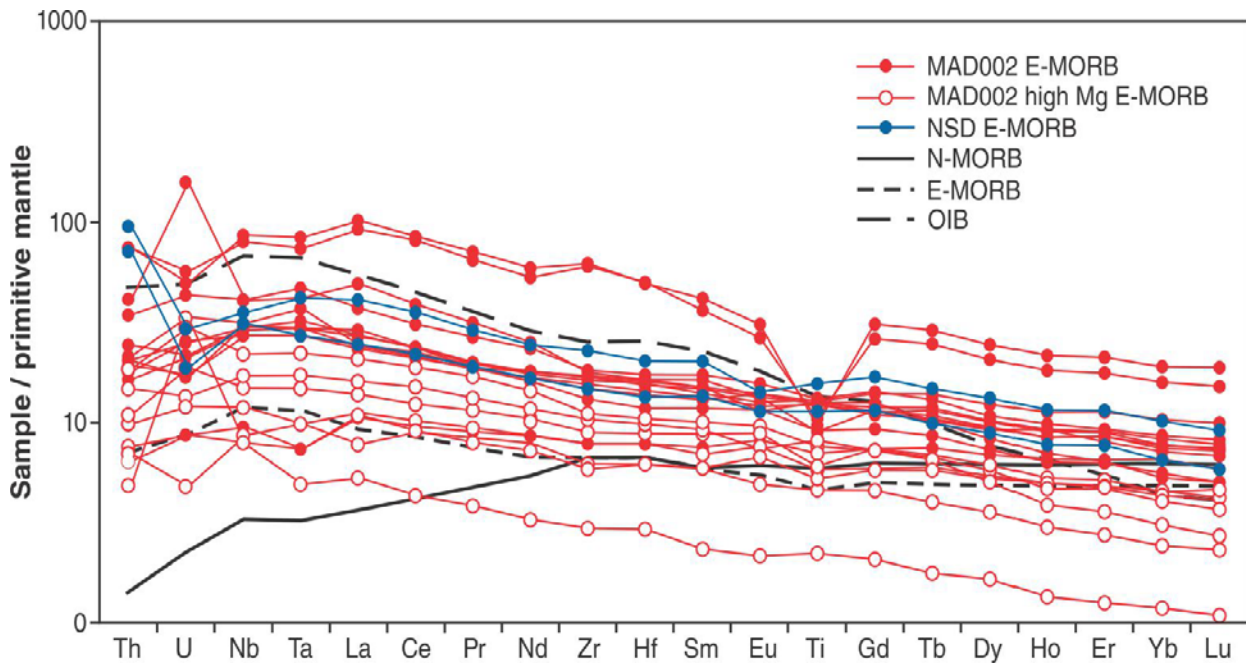


Pre- c. 1200 Ma tectonic evolution

- <1950 – 1600 Ma history of the AFO is one of an ocean-continent transition
- The ocean created at this stage is the c. 1950 Ma mafic crust identified in our isotopic data – i.e. the crust that became the Madura and Musgrave Province and Forrest Zone
- Partially consumed with convergence creating multiple oceanic arcs from c. 1720 – 1600 Ma (FZ) and at c. 1400 Ma (MP)

Basalts constrained between 1950 and 1400 Ma



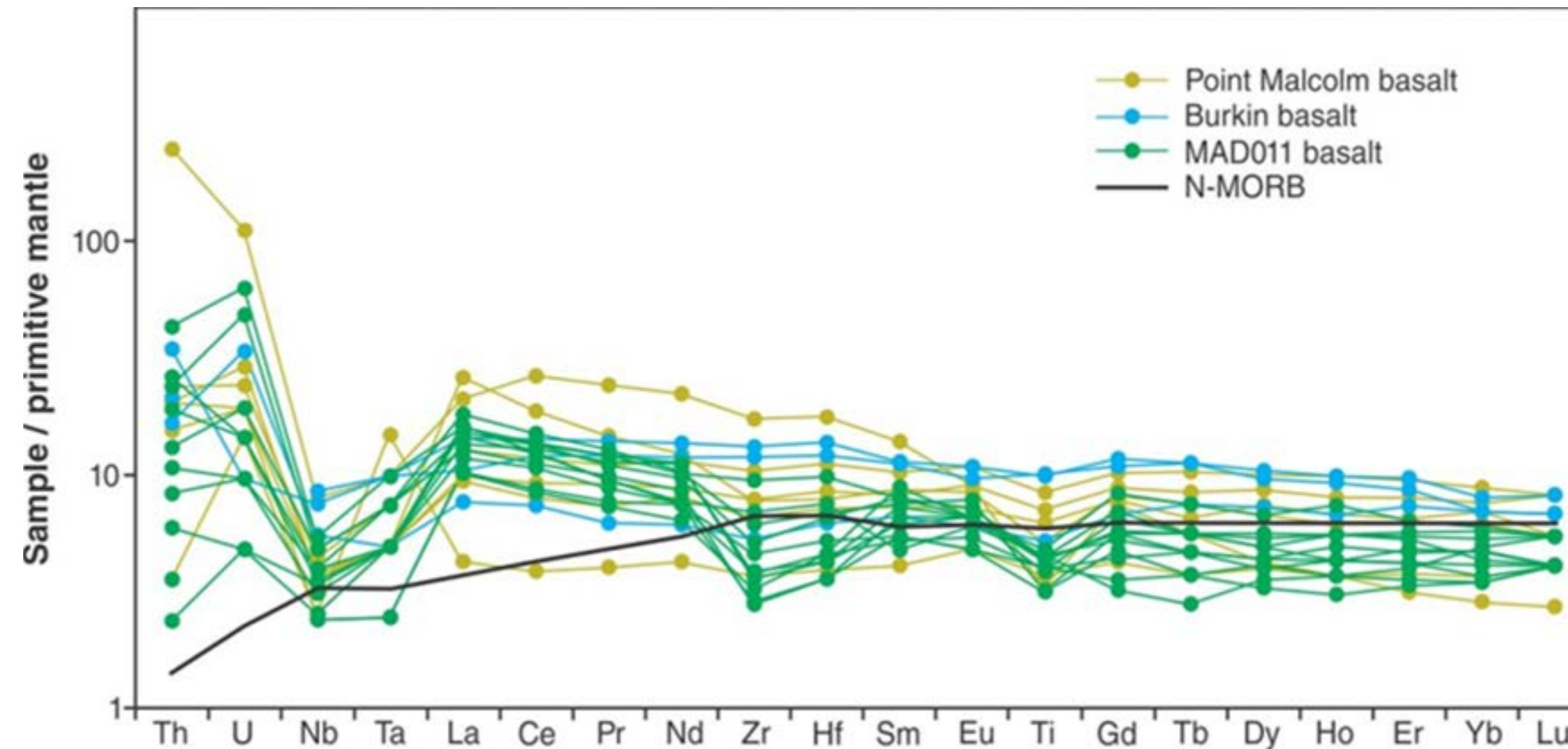


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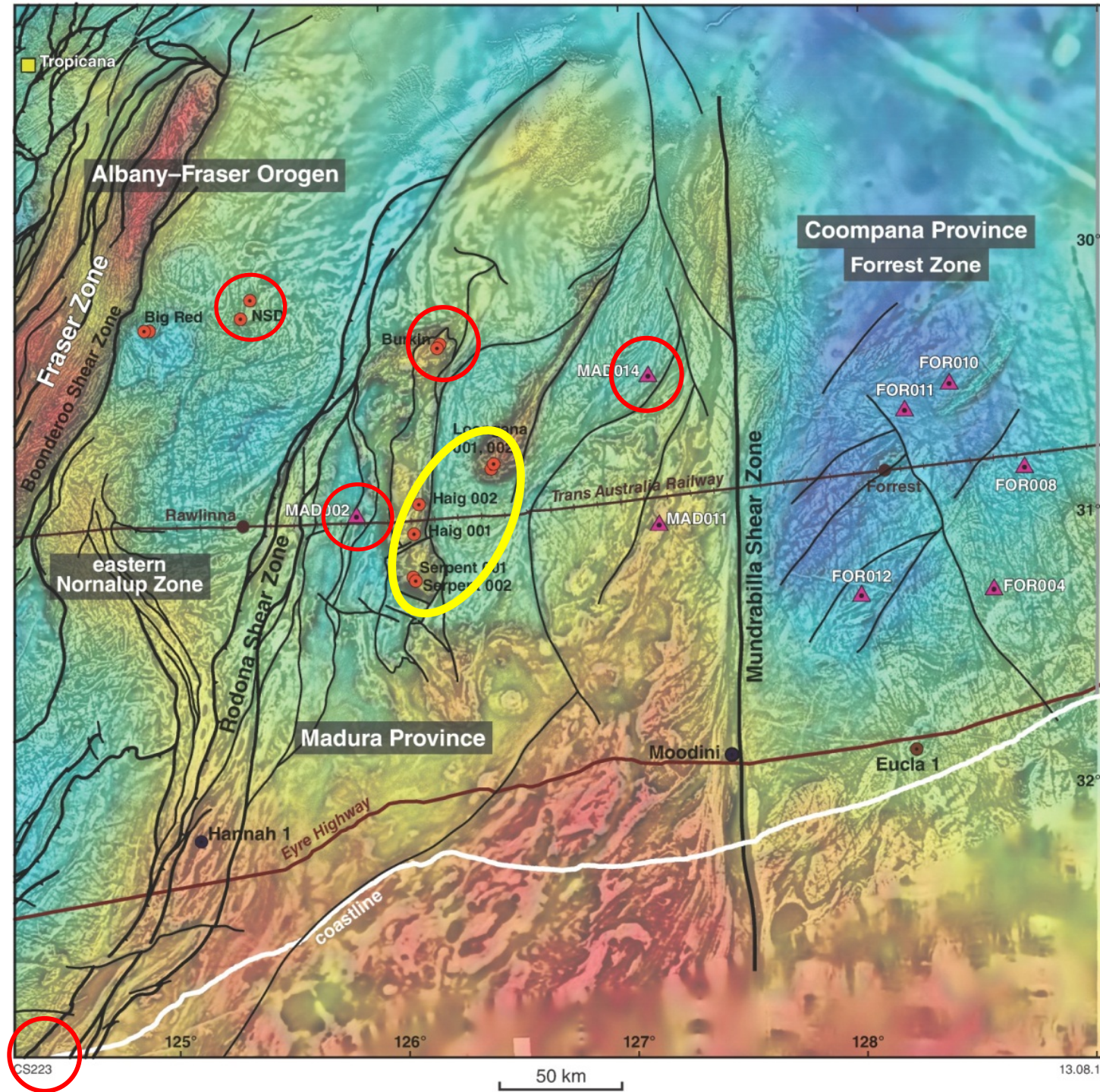
- Primitive ($Mg^\#$ up to 75, very high Cr and Ni)
- Strongly enriched, with no negative Nb anomaly
- Trace element patterns similar to E-MORB
- High Ti/Yb ratios reflecting a garnet-residual mantle source

Reflects asthenospheric upwelling and are an expected result of continental rift → OCT.

2



- Relatively flat to weakly enriched normalized trace element patterns.
- Trends for HFSE and Yb suggest a source very similar to N-MORB source.

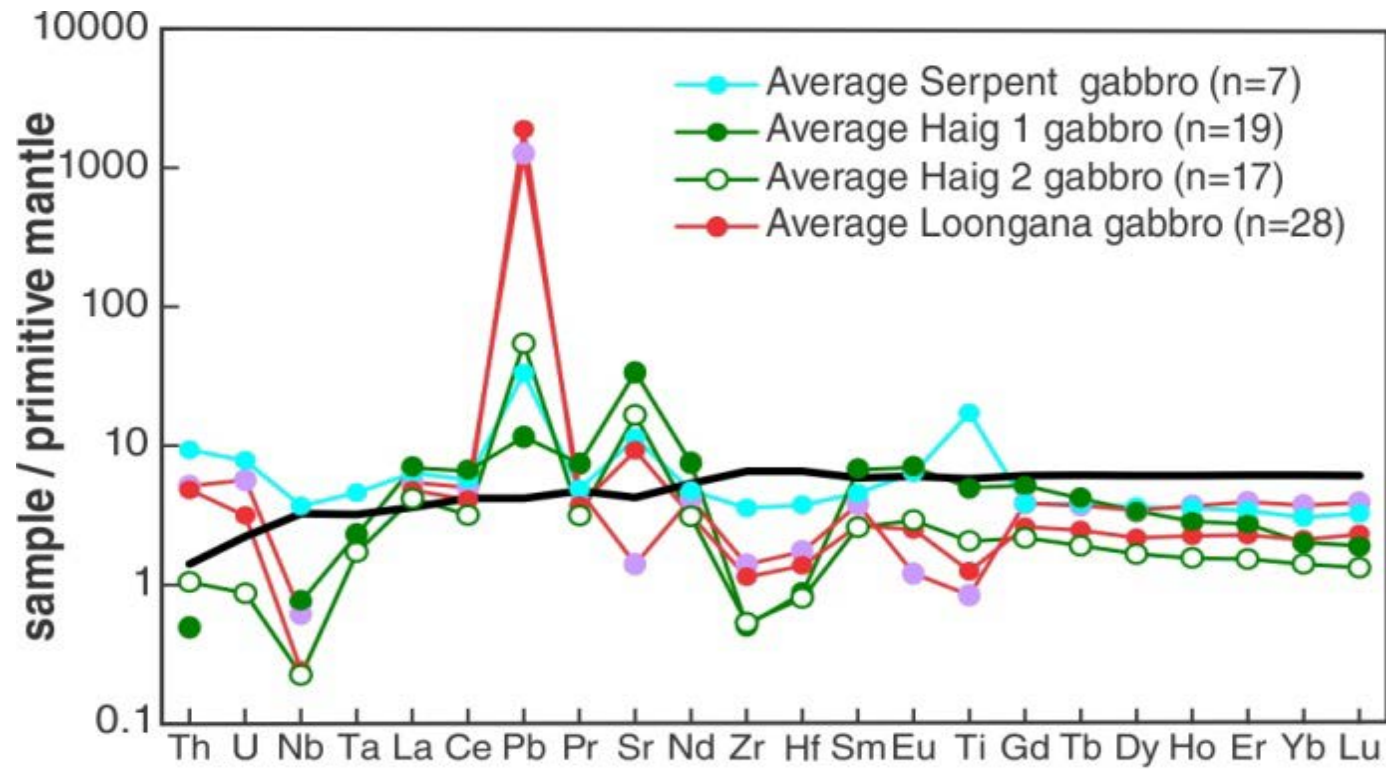


Gabbros

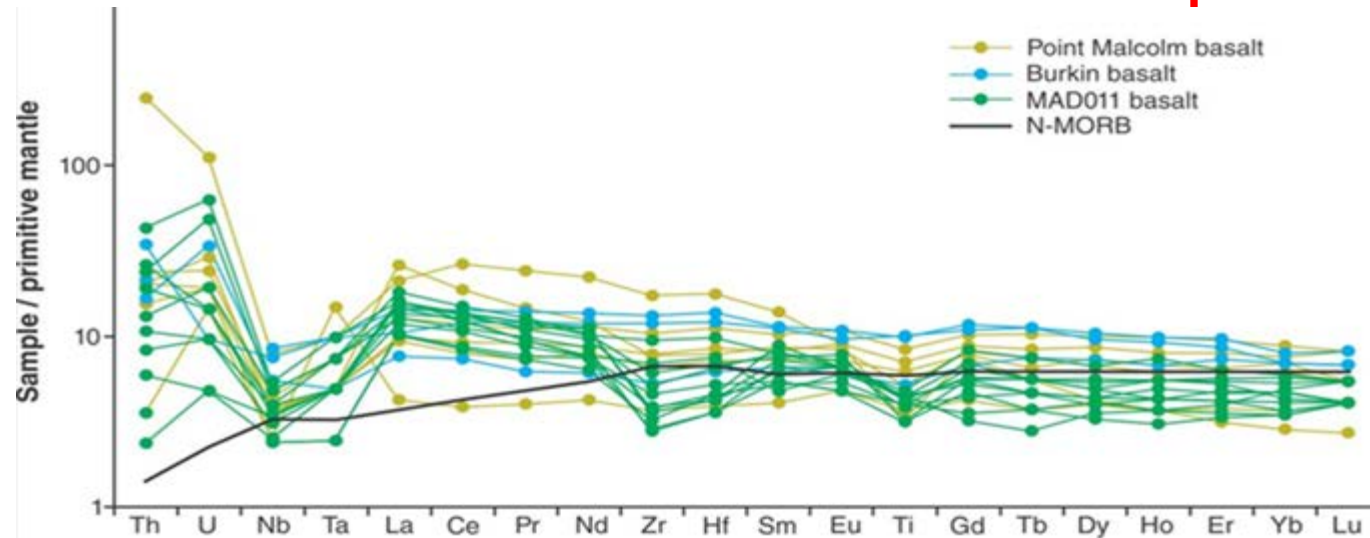
Loongana

Haig

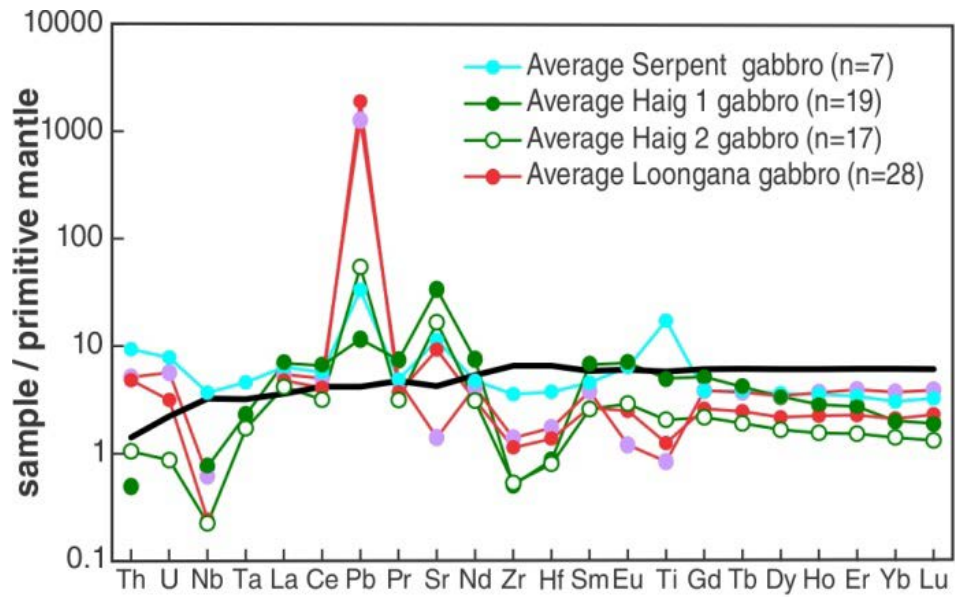
Serpent



- Broadly low-K tholeiitic
- Same flat to weakly enriched normalised trace element patterns, with negative Nb and Zr, as the **second (eastern) group of basalts** – but with some positive ‘spikes’ reflecting cumulate minerals.
- Also in common with those basalts, trends in HFSE (Nb, Zr, Ti) and Yb broadly parallel N-MORB and reflect a similarly depleted mantle source (N-MORB source)



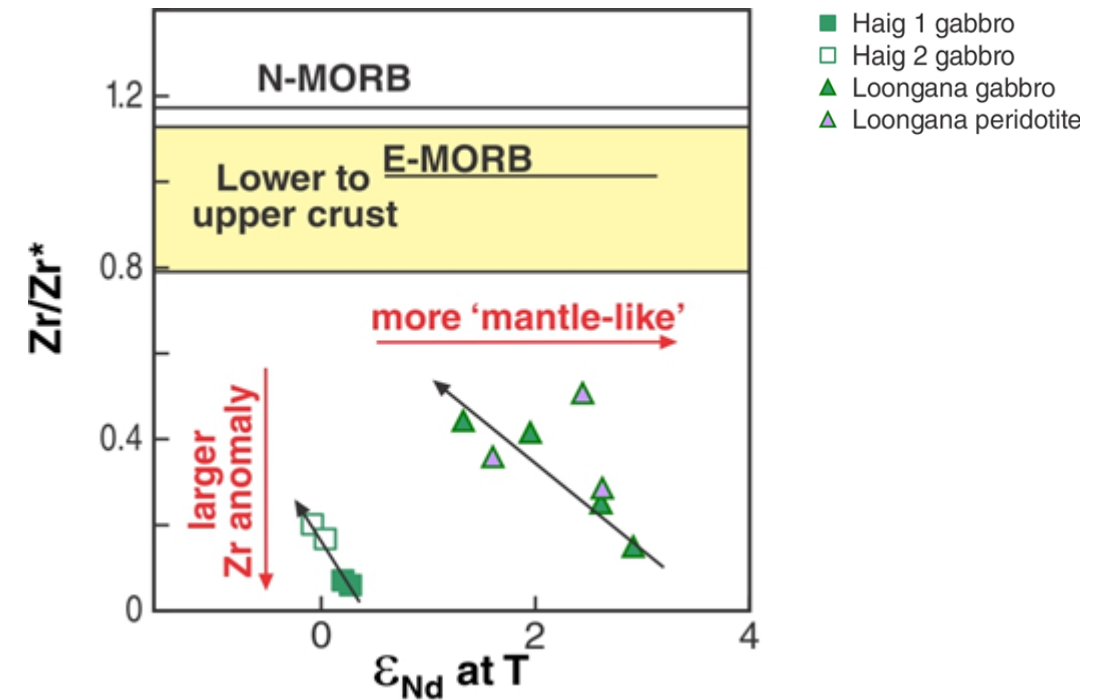
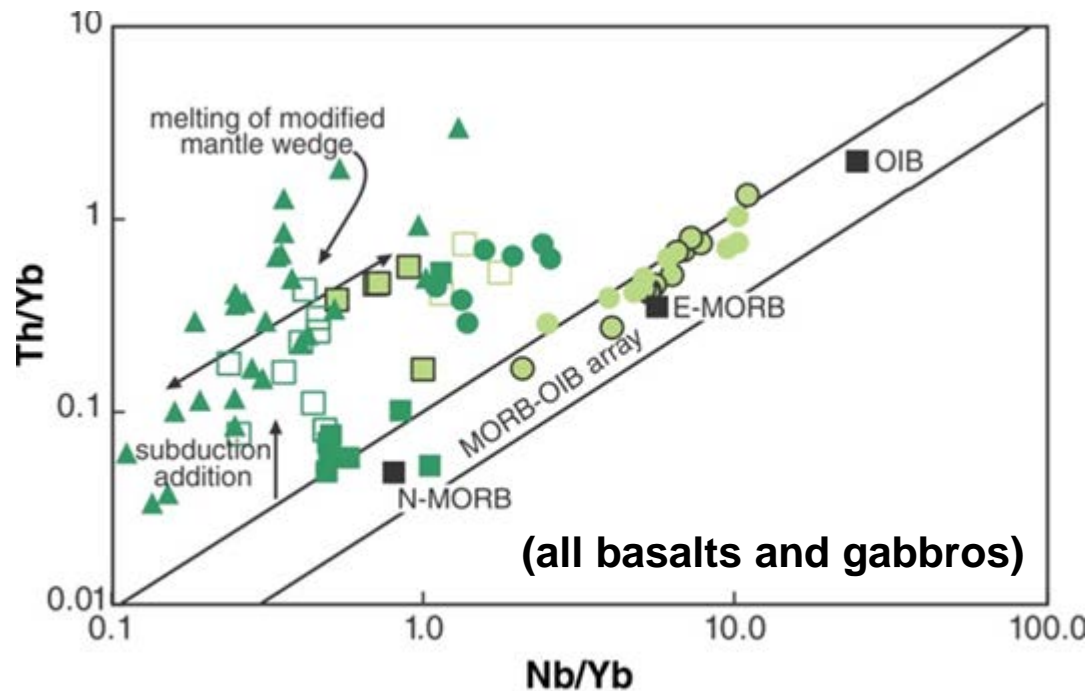
Source and petrogenesis probably very similar to the basalts



Enriched patterns reflect either a subduction-enriched mantle source (i.e. subduction) OR simply crustal contamination of a mantle-derived melt.

NEED to demonstrate which applies!!!

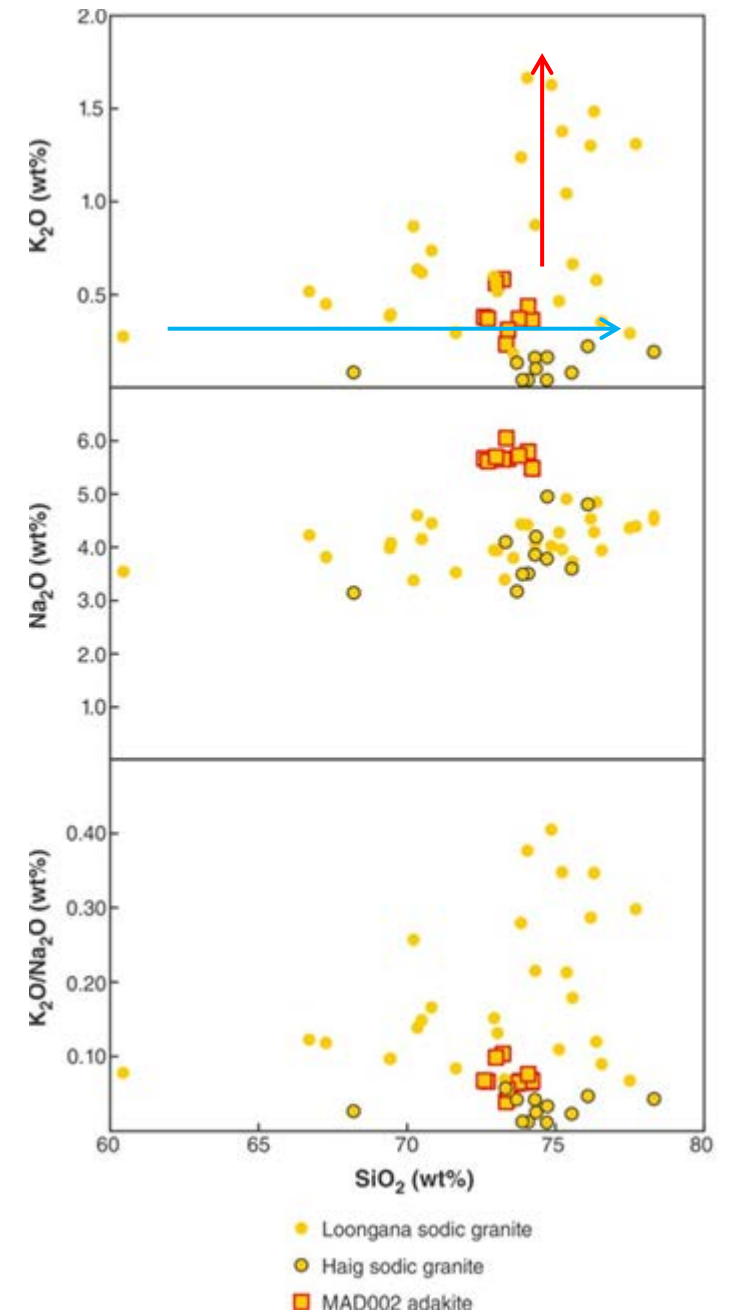
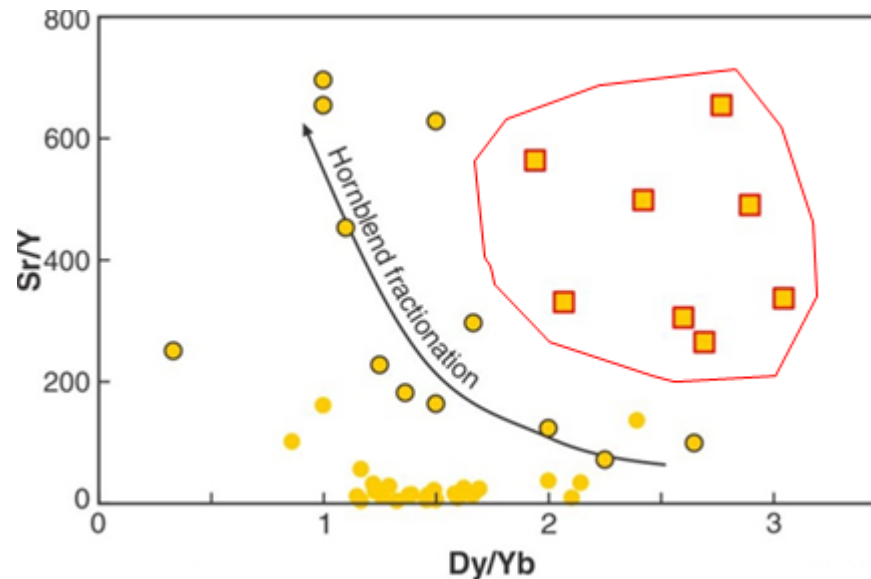
Isotopically primitive terrain gives us one clue – but various trace element arguments also provide very strong support for a subduction setting.



Gabbros and the E-MORB in MAD002 are intruded by c. 1400 Ma sodic granites

Two types recognised

- 1) 'Oceanic' plagiogranites,
 - a) some showing a purely 'oceanic' (i.e. mafic) source (→)
 - b) some showing slight K-enrichment (source incorporates minor hyper extended AFO crust)(→)
- 2) Adakite – intruding E-MORB in MAD002. Very high Sr/Y and La/Yb ratios and extreme depletions in HREE reflecting high-pressure melting of a subducted oceanic slab.

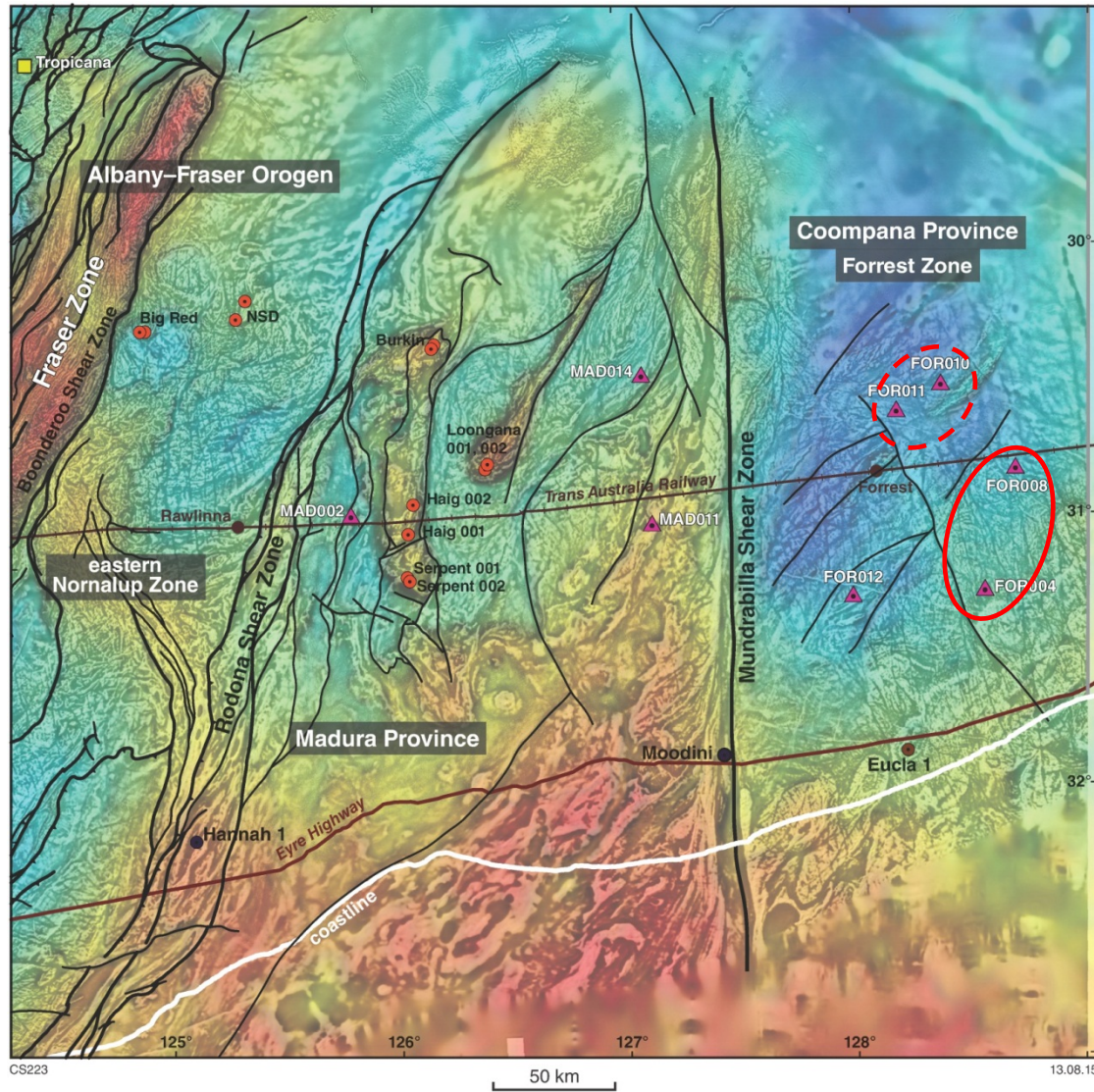


Madura Province (pre c. 1200 Ma)

- **Early E-MORB in the west – remnants of the early ocean-continent transition phase.**
- **Widespread basalt and c. 1400 Ma gabbro with the signature of a N-MORB source that was weakly subduction-modified - likely reflects oceanic arc or fore-arc.**
- **Plagiogranites form part of the oceanic arc sequence and include types typically found in ophiolites and most clearly evolved from a source free of significant 'continental' material**
- **1400 Ma adakites intrude much older E-MORB, indicate hot subduction of oceanic lithosphere, and likely formed to the east of an east-dipping slab.**

Forrest Zone c. 1610 Ma (c. 1720 and 1670 Ma inheritance)

Subduction-related magmatism: arc crust



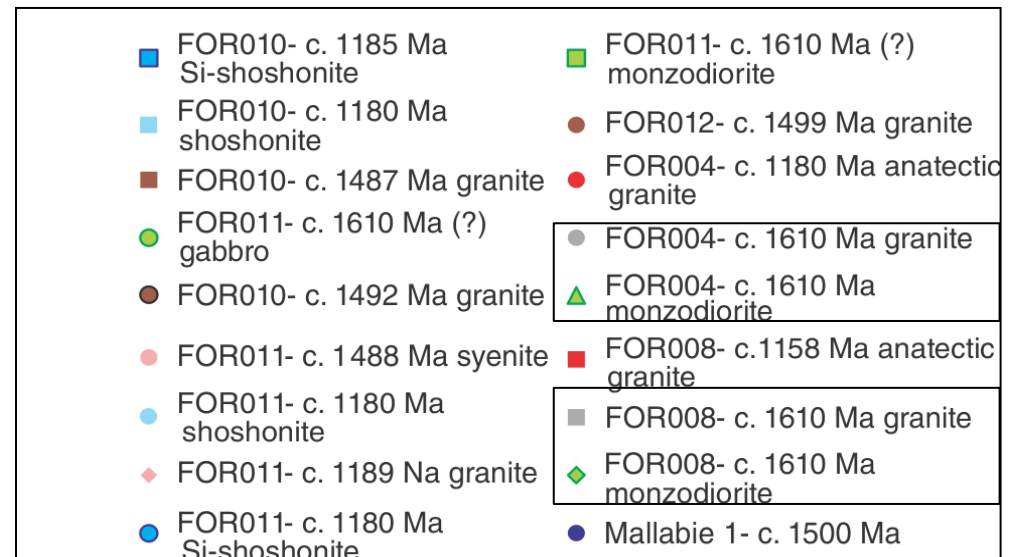
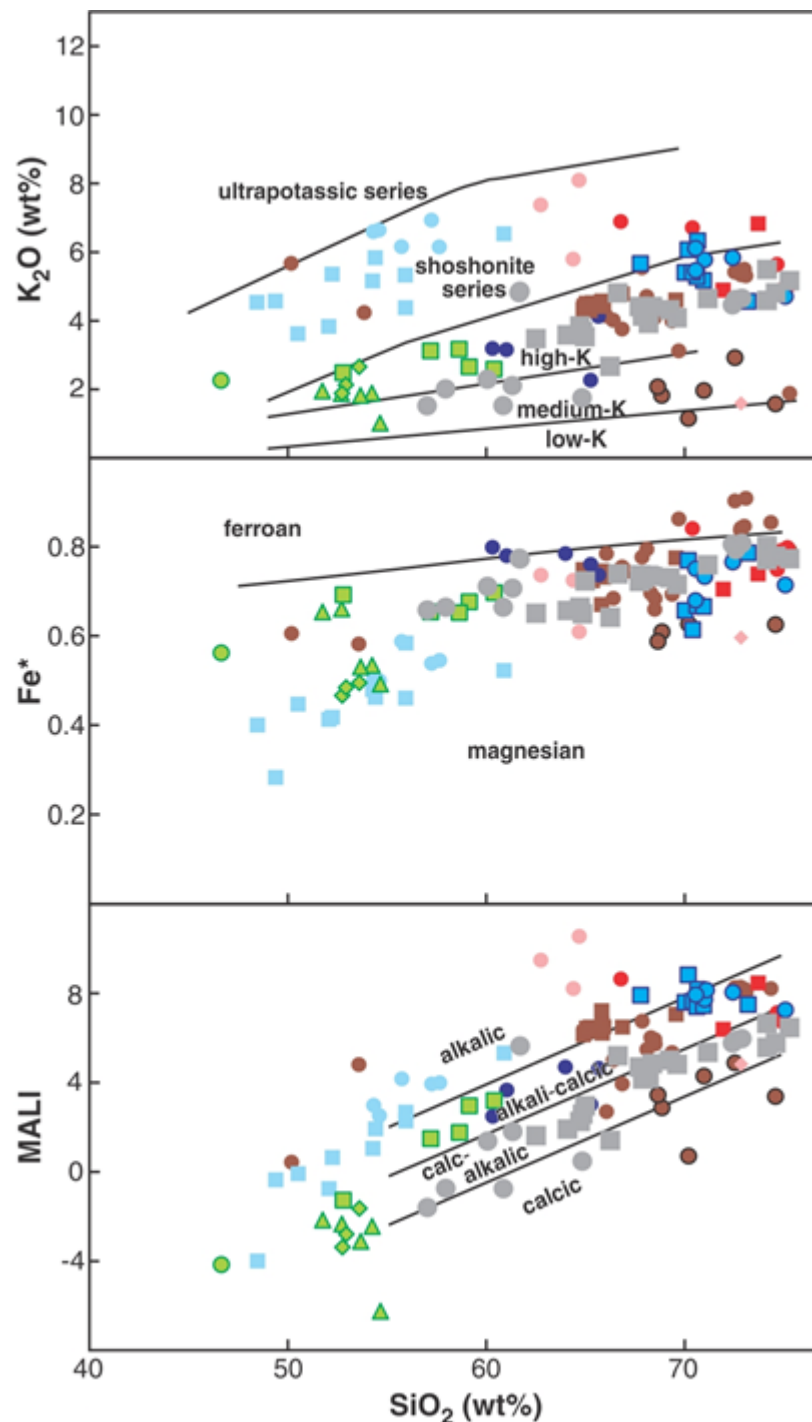
Toolgana
Supersuite

FOR004
FOR008

Undated but
compositionally the
same
(FOR011)
(FOR010)

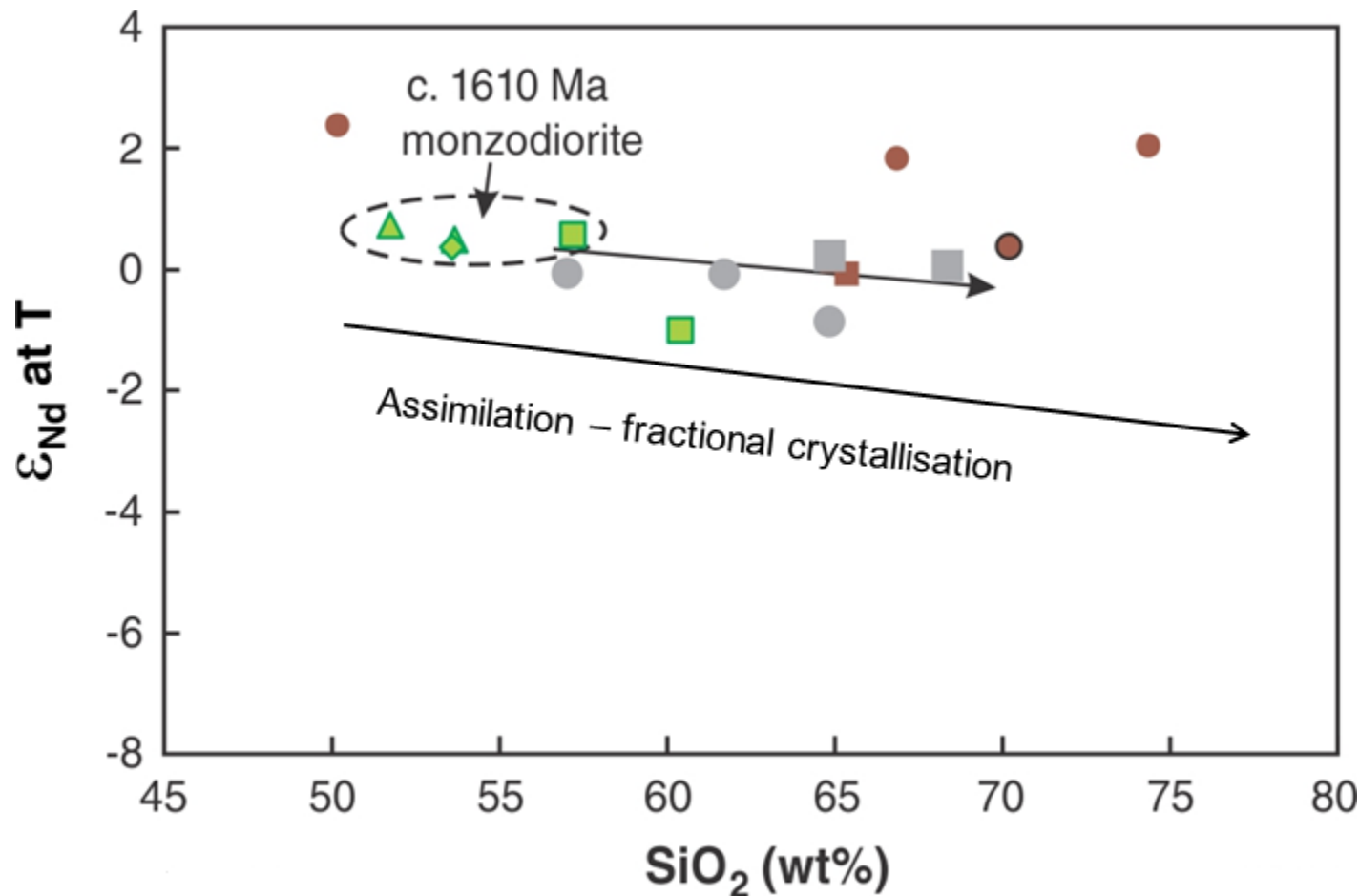
Toolgana Supersuite

- Granites are magnesian and mainly calc-alkalic
- Mainly high-K
- Associated with abundant mafic equivalents



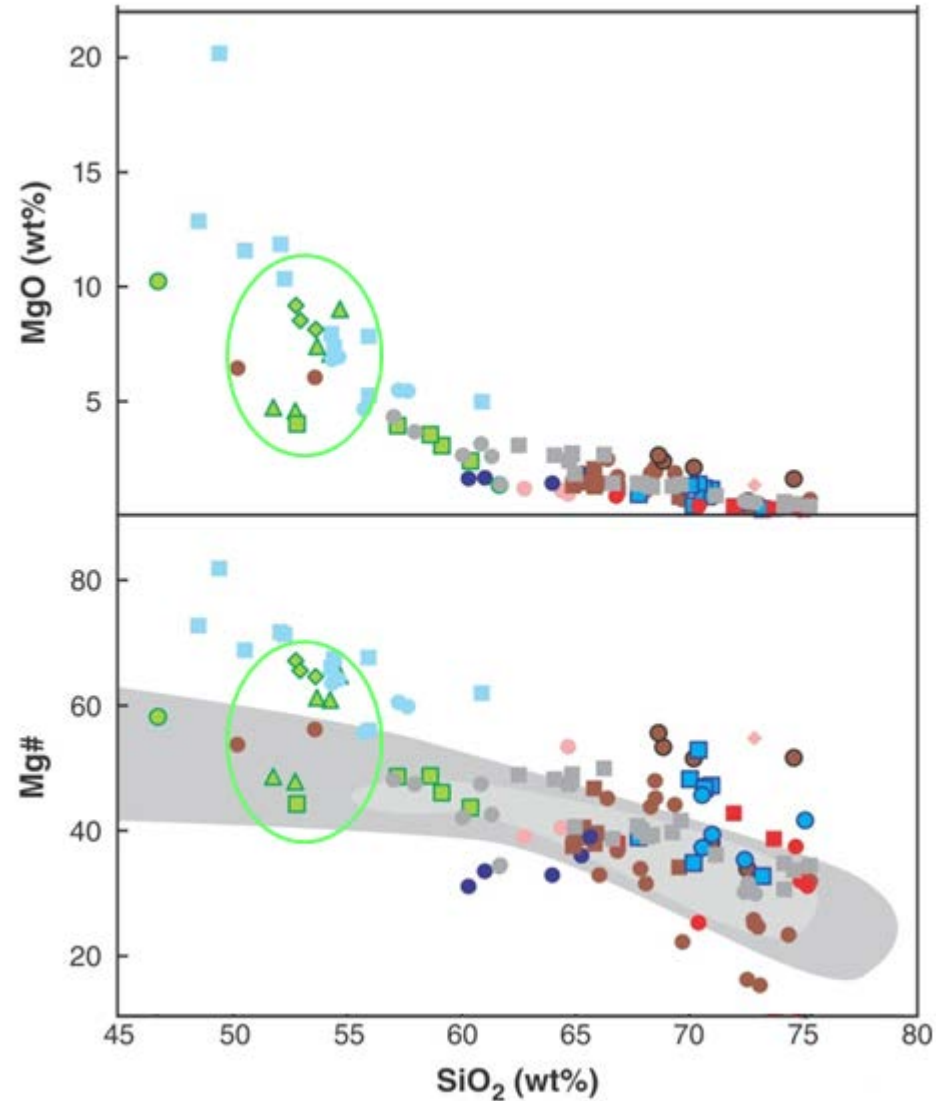
Toolgana Supersuite

Nd isotopic composition on the granites become slightly less radiogenic over a large increase in SiO_2 , suggesting AFC type processes involving only very minor progressive contamination (? the inherited 1724 – 1671 Ma component) and/or an isotopically similar contaminant.



Monzodiorite and granite form co-genetic suites that have intruded c. 1720-1670 Ma crust that is compositionally similar – recycling of slightly earlier arc crust - **Composite arcs!!**

Toolgana Supersuite - monzodiorite

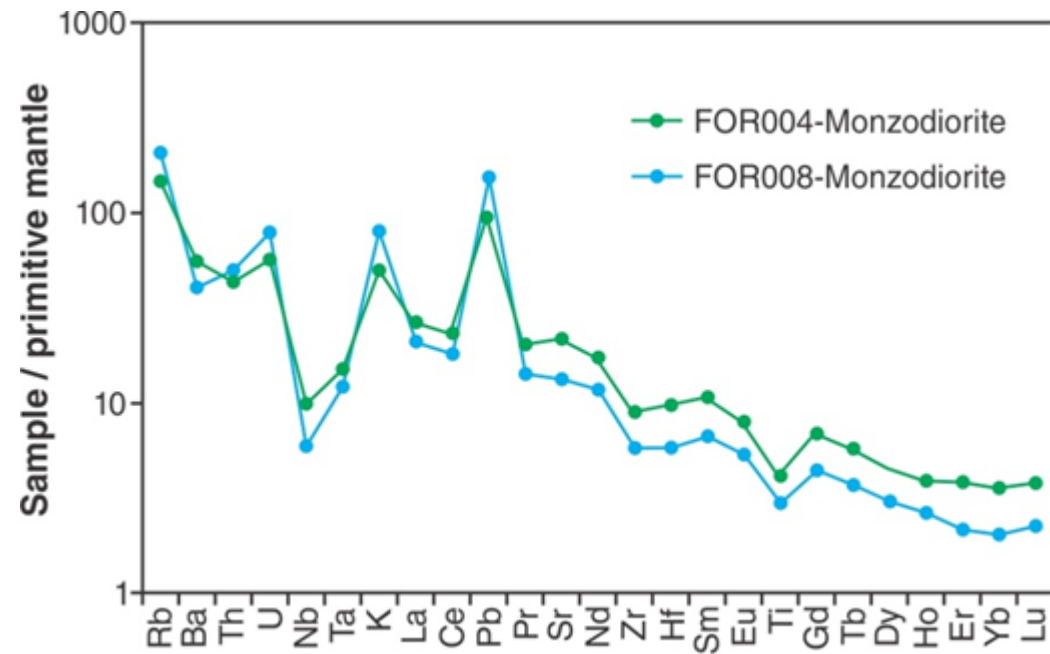


The most primitive samples have:

SiO₂ ~ 52 wt%
MgO > 9.0 wt%
Mg > 65
Ni up to 187 ppm

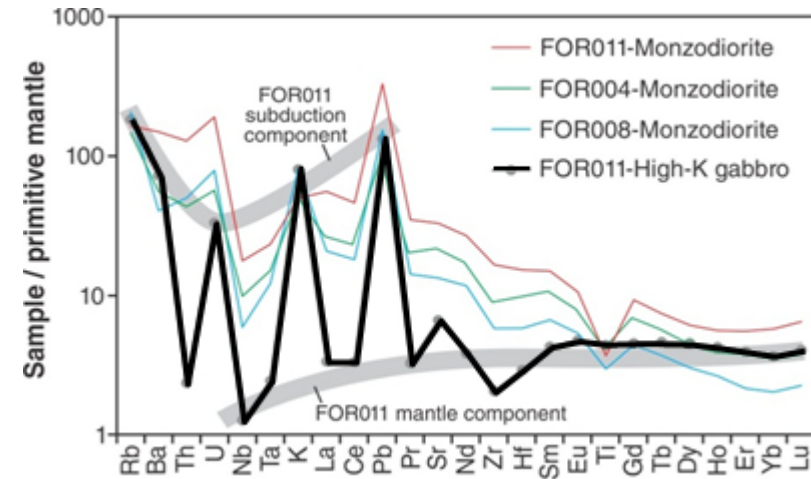
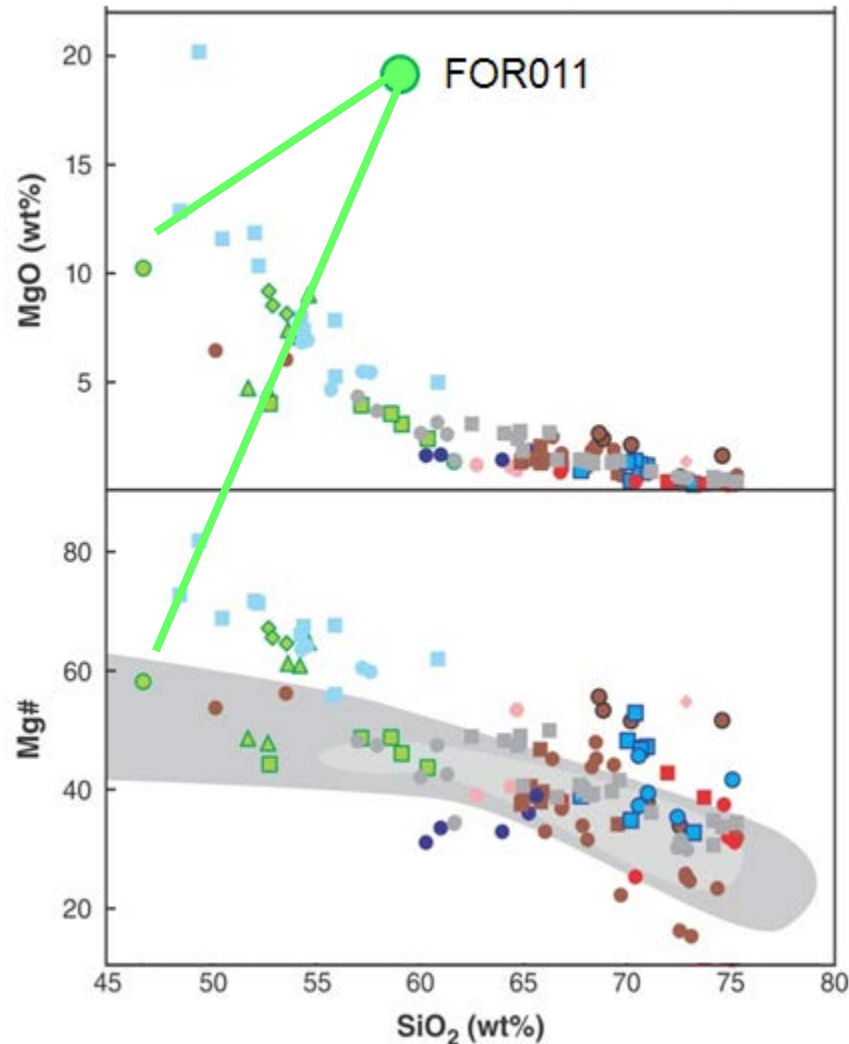
Rules out significant
fractional
crystallisation

But are also significantly enriched in highly
incompatible trace elements, with high
LILE/HFSE ratios (i.e. subduction-like patterns)



Toogana Supersuite

K-rich gabbro (2.26 wt% K_2O at 46.63 wt% SiO_2) also occurs in FOR011.



Depleted mantle source

Significant enrichment in fluid-mobile trace elements

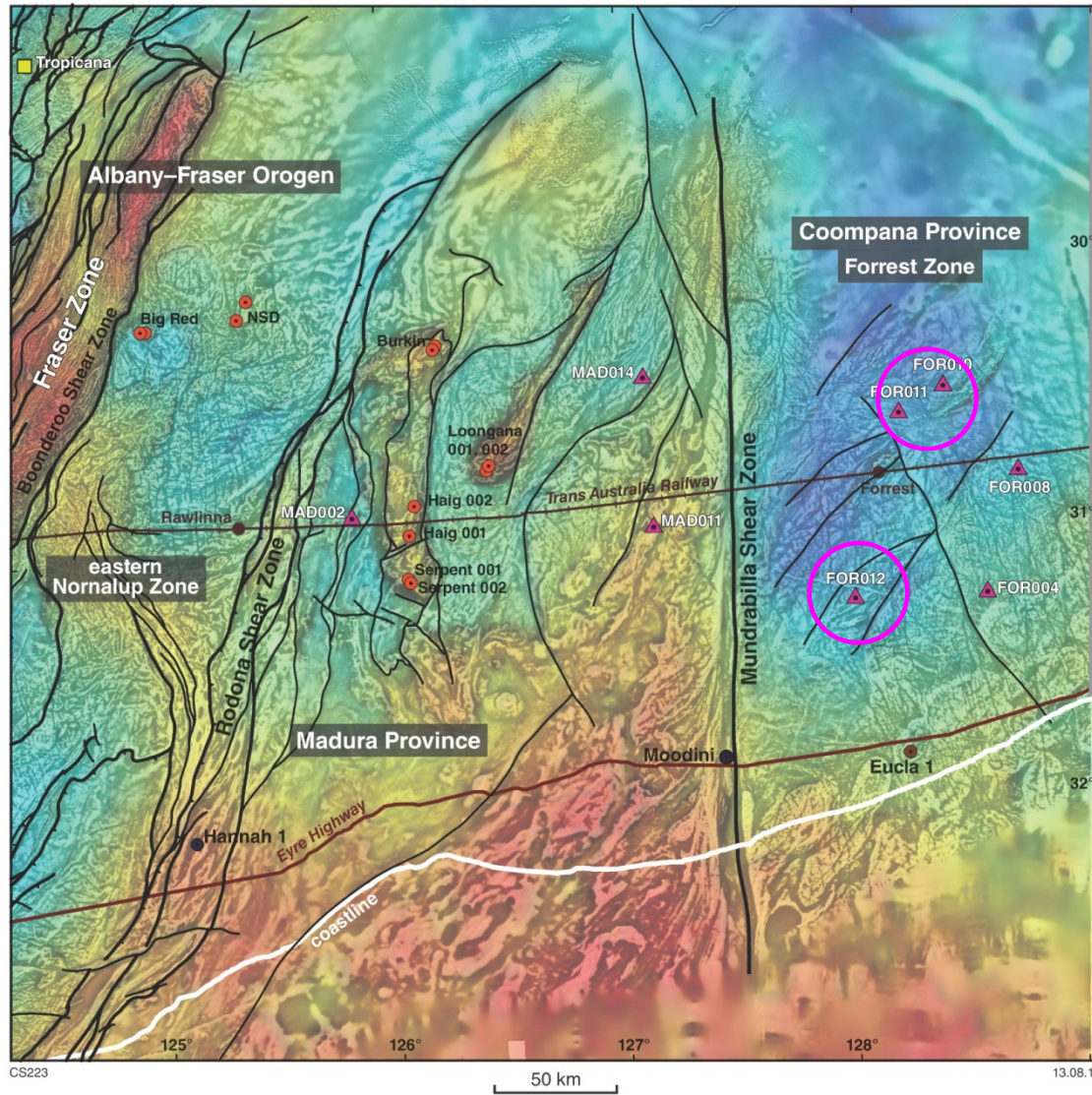
NO crustal addition

Enrichment pattern that resembles that in other c. 1610 Ma rocks

Fore arc or back arc tholeiite; and added confidence that the Toogana SS is subduction-related

Forrest Zone c. 1500 Ma

Extension (rift) related granites



Undawidgi Supersuite

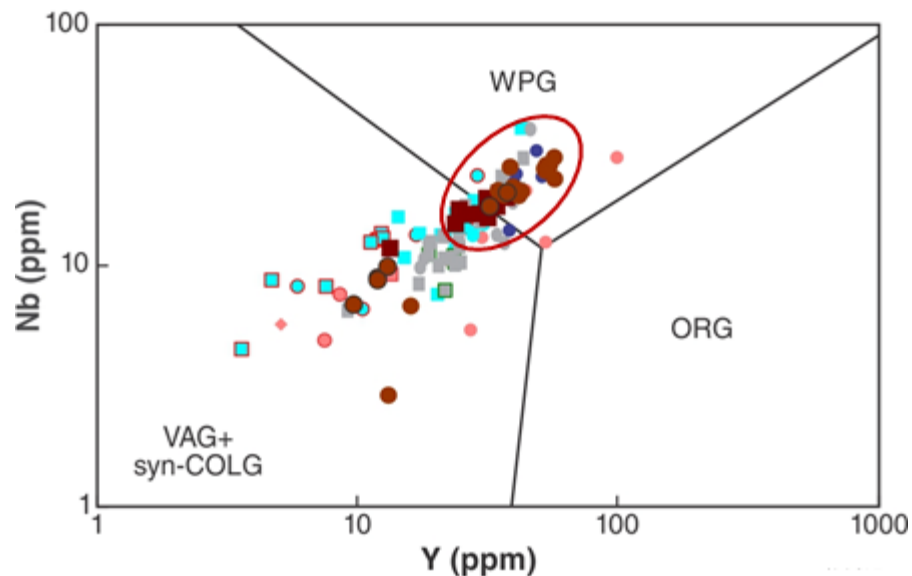
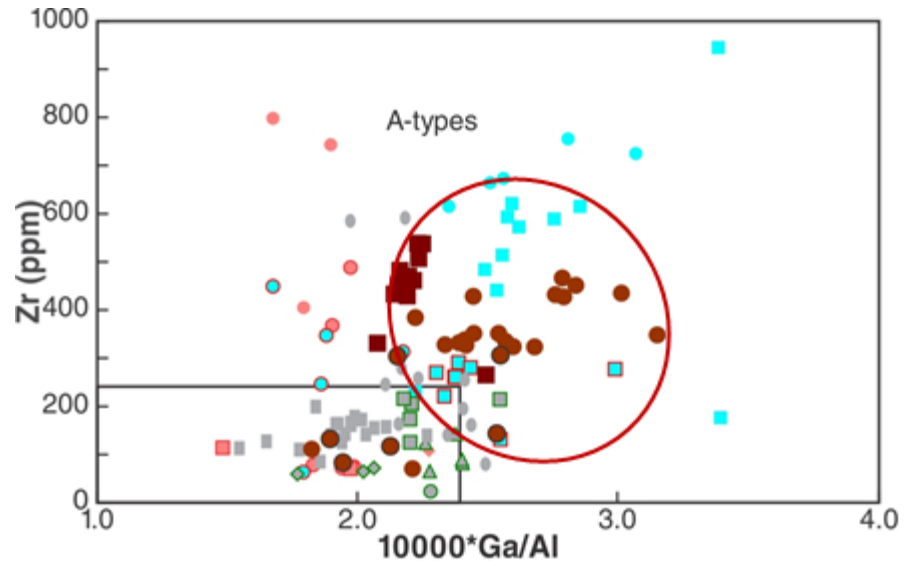
FOR012

FOR010

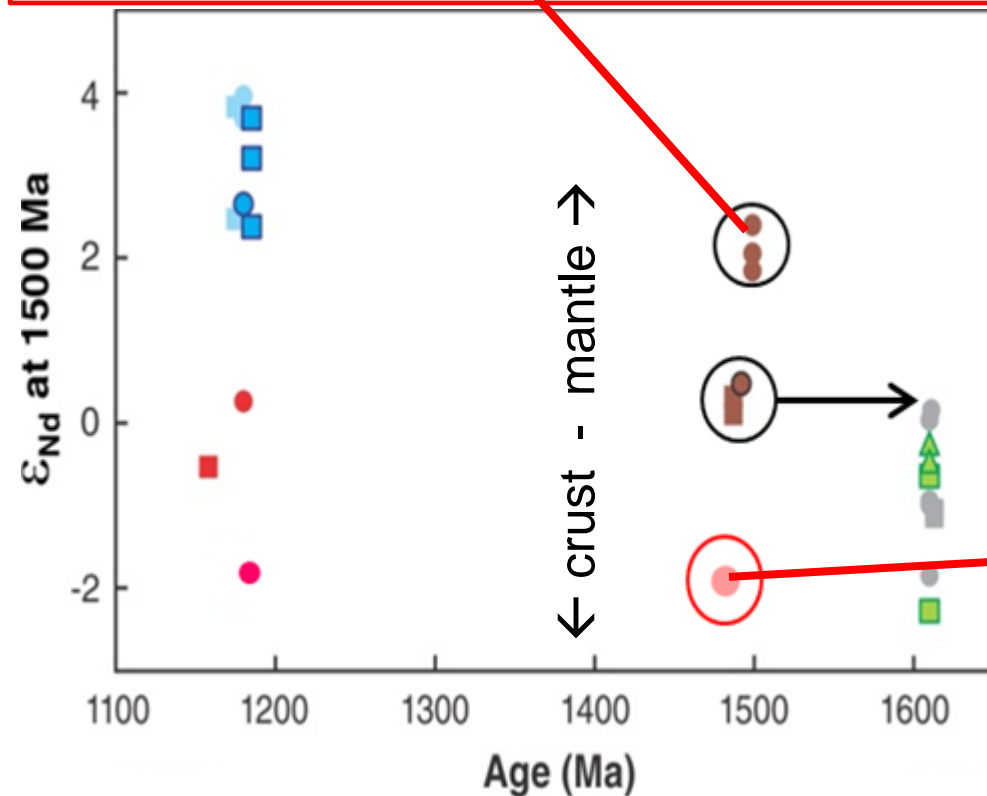
FOR011

High-K Undawidgi Supersuite

Show enrichments in TiO_2 , P_2O_5 , HFSE, Ga, Zn, REE typical of hot and dry A-type magmas.



High-K granites ! addition of a more radiogenic component (e.g. mantle) — and if this was upwelled asthenospheric mantle this would additionally be consistent with their more ferroan compositions



Syenite
? Remelting of strongly alkali-metasomatised Toolgana crust

Forrest Zone (pre c. 1200 Ma)

- **Toolgana Supersuite c. 1600 Ma subduction-related (super)suite. Little isotopic evidence for the type of 'continental' input we would expect to see at a continental margin – so probably largely oceanic. History of arc magmatism and recycling spans at least 1720 – 1600 Ma (earlier than the preserved subduction record in the Madura Province)**
- **Undawidgi Supersuite c. 1500 Ma Range of rocks each with compositional features that distinguish them from earlier arc-related rocks (source). Influence of hot, dry, asthenospheric input likely in an extensional environment (rift, arc extension)**

**late (1220 - 1150 Ma) reworking of entire region,
including the cratonic margins**

The Maralinga Event

Albany-Fraser Orogeny Stage II

Musgrave Orogeny

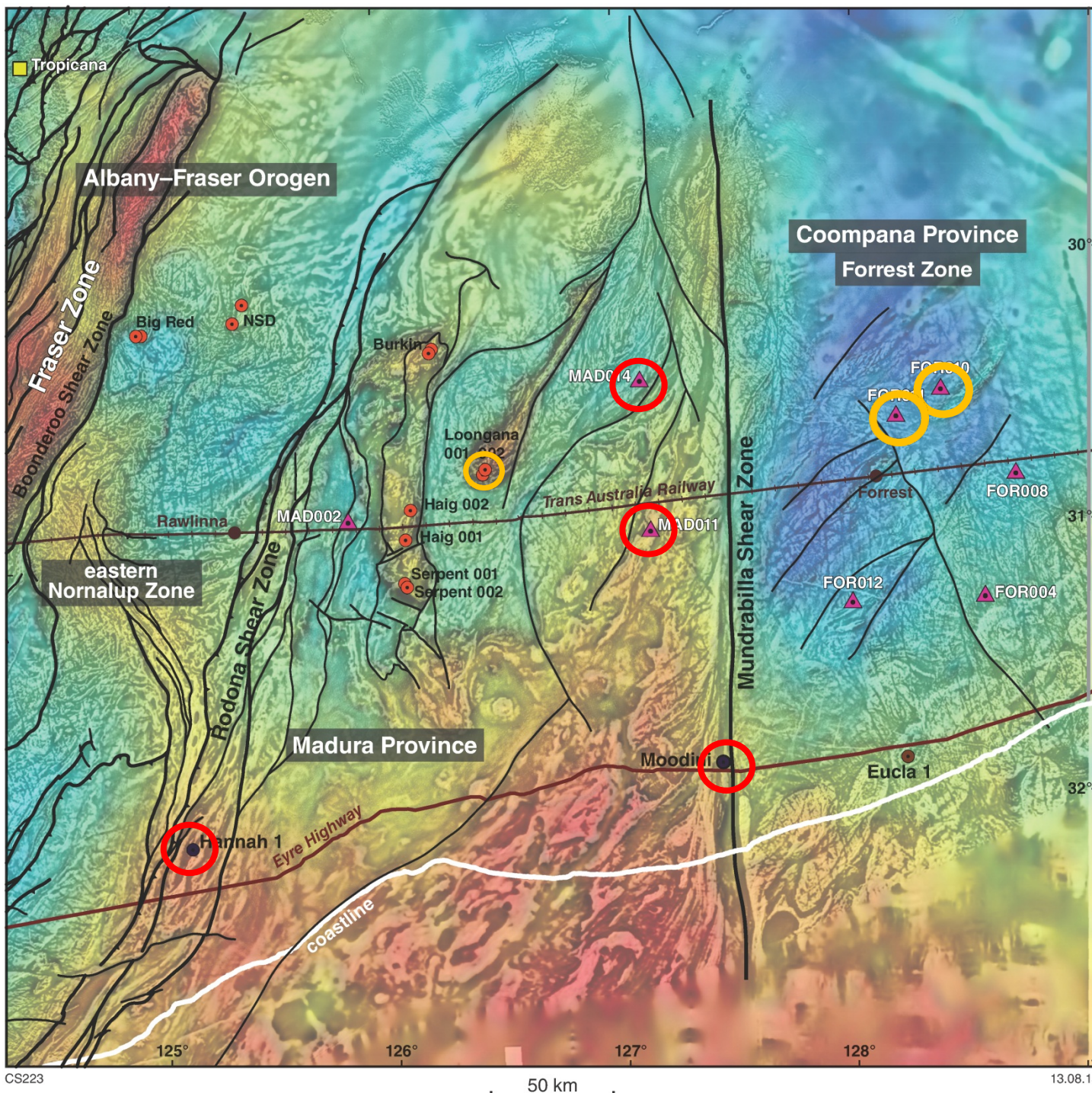
**Great central
Australian meltdown**

Madura Province and Forrest Zone (1190 Ma onwards)

- **Our record of oceanic plate consumption and arc formation ends at c. 1400 Ma.**
- **Final stages of convergence continues as oceanic terrains were thrust over the AFO (not sure what happened on the other ?"Gawler" side of the ocean).**
- **Very little activity is recorded between c. 1290 and 1220 Ma anywhere in the Madura, Forrest, Musgrave or Albany-Fraser regions.**
- **At c. 1200 Ma however, we begin the most extraordinary regional magmatic or thermal event which fundamentally changes the geochemical composition – but NOT the isotopic heritage – of large portions of the entire region.**

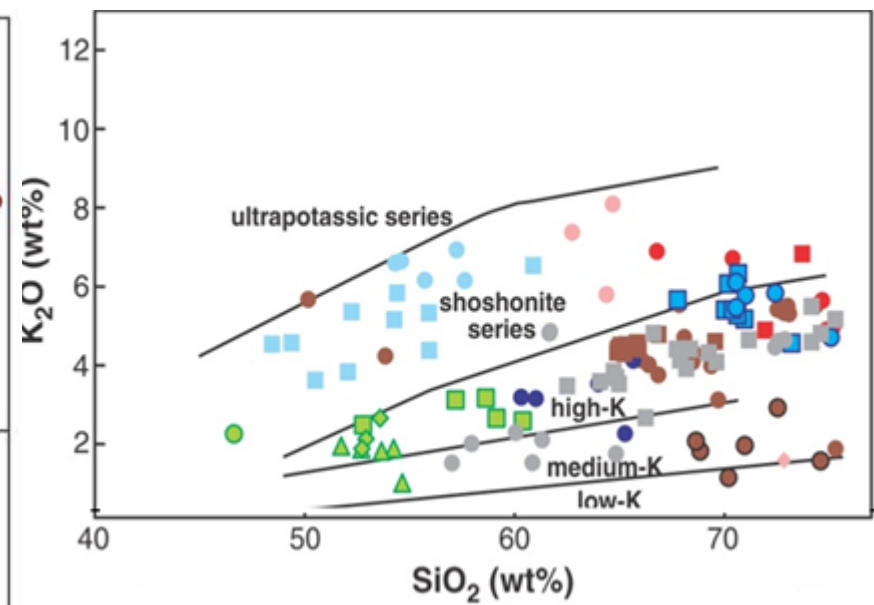
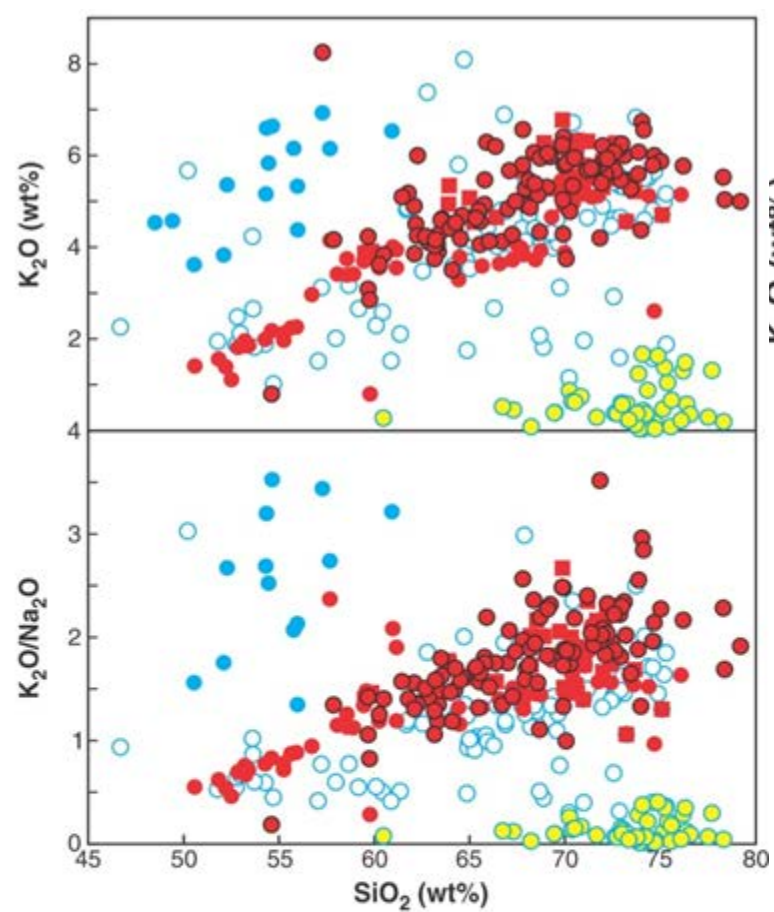
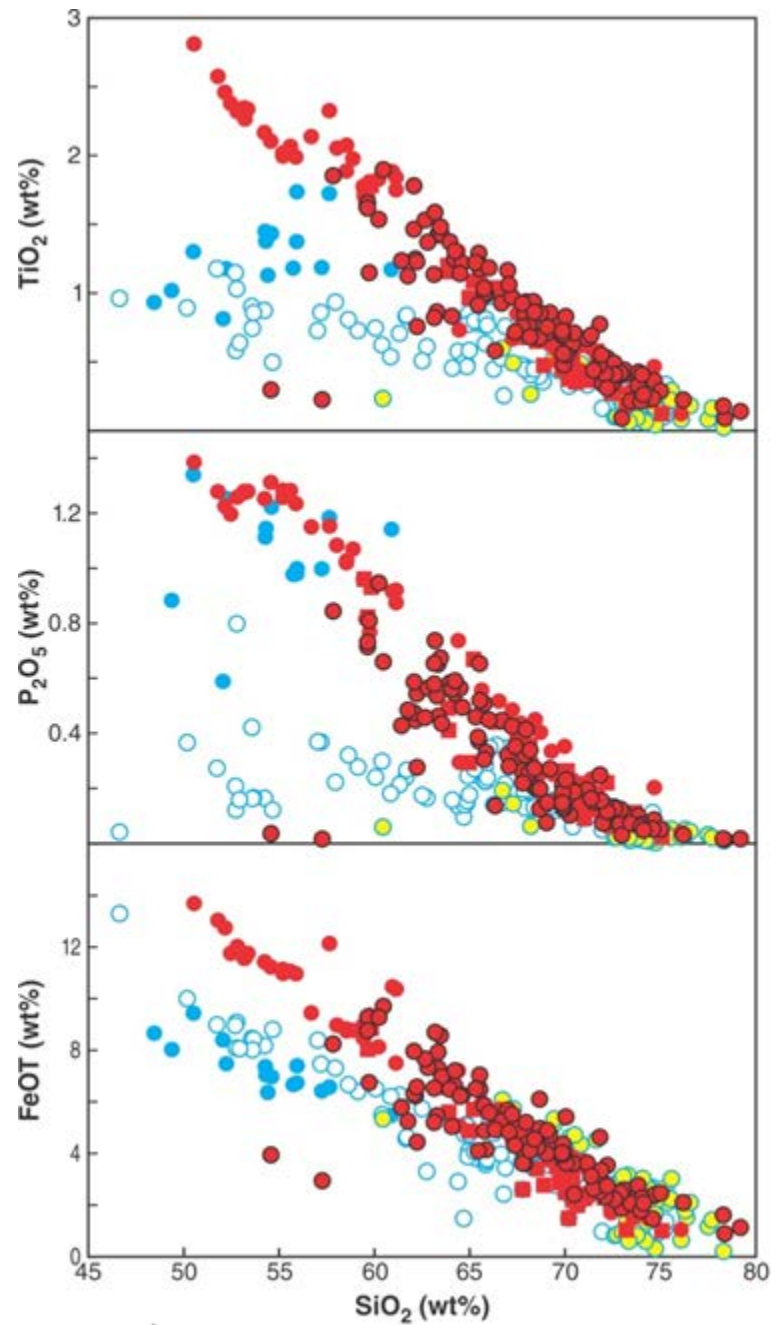
Moodini Supersuite - Two strongly contrasting styles of magmatism

- 1) Madura Province c. 1180 – 1125 Ma**
High-KFe (ferrodiorite to A-type granite) magmatism. Regionally voluminous, mainly in the Madura and Musgrave Provinces but also in the AFO; high magnetic signature. (*Not sampled in the Forrest Zone but aeromagnetic traces suggest their presence.)
- 2) Forrest Zone c. 1185 – 1150 Ma**
High-KMg (shoshonite and shoshonitic granite) magmatism – mainly in the Forrest Zone. (*Voluminous in the Forrest Zone but only one occurrence – at Loongana – of thin dykes in the Madura Province.)



High-KFe series: high-T melting of extended lower crust

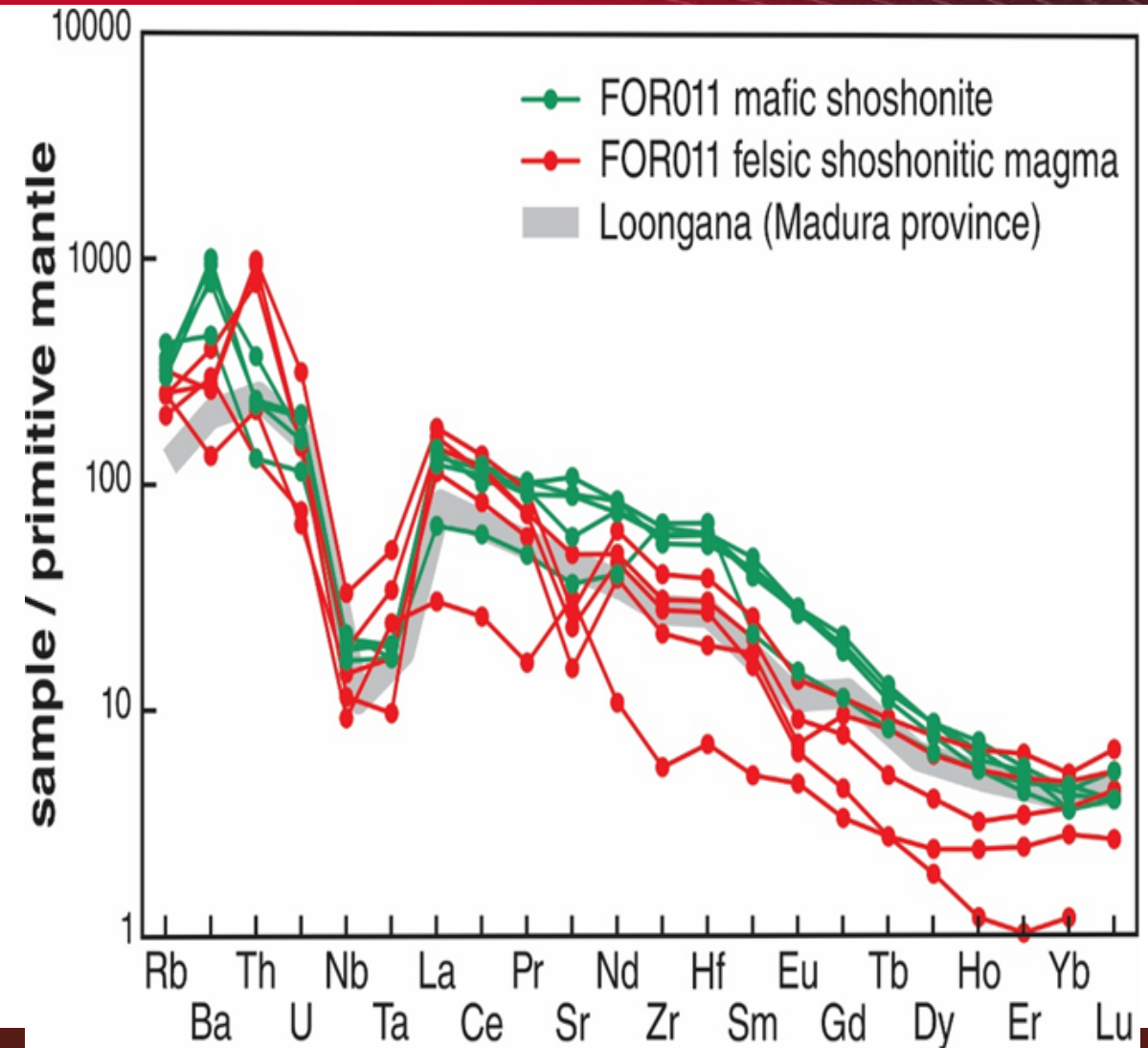
- Include orthopyroxene- and pigeonite-bearing monzodiorite and monzonite (jontunites) to biotite-monzogranite.
- Geochemically specialised – high K, Fe, Ti, P, HFSE, REE, Th, U – A-type (or ‘Charnockite type’)
- Intrusive temperatures often $> 900^{\circ}\text{C}$ (mafic member $>1000^{\circ}\text{C}$).
- Geochemical and isotopic characteristics suggesting melting, under very high temperatures and dry and reduced conditions, of refractory lower-crust with significant material and thermal contribution from asthenospheric mantle.



High-KMg series – melting of subduction metasomatised lithosphere



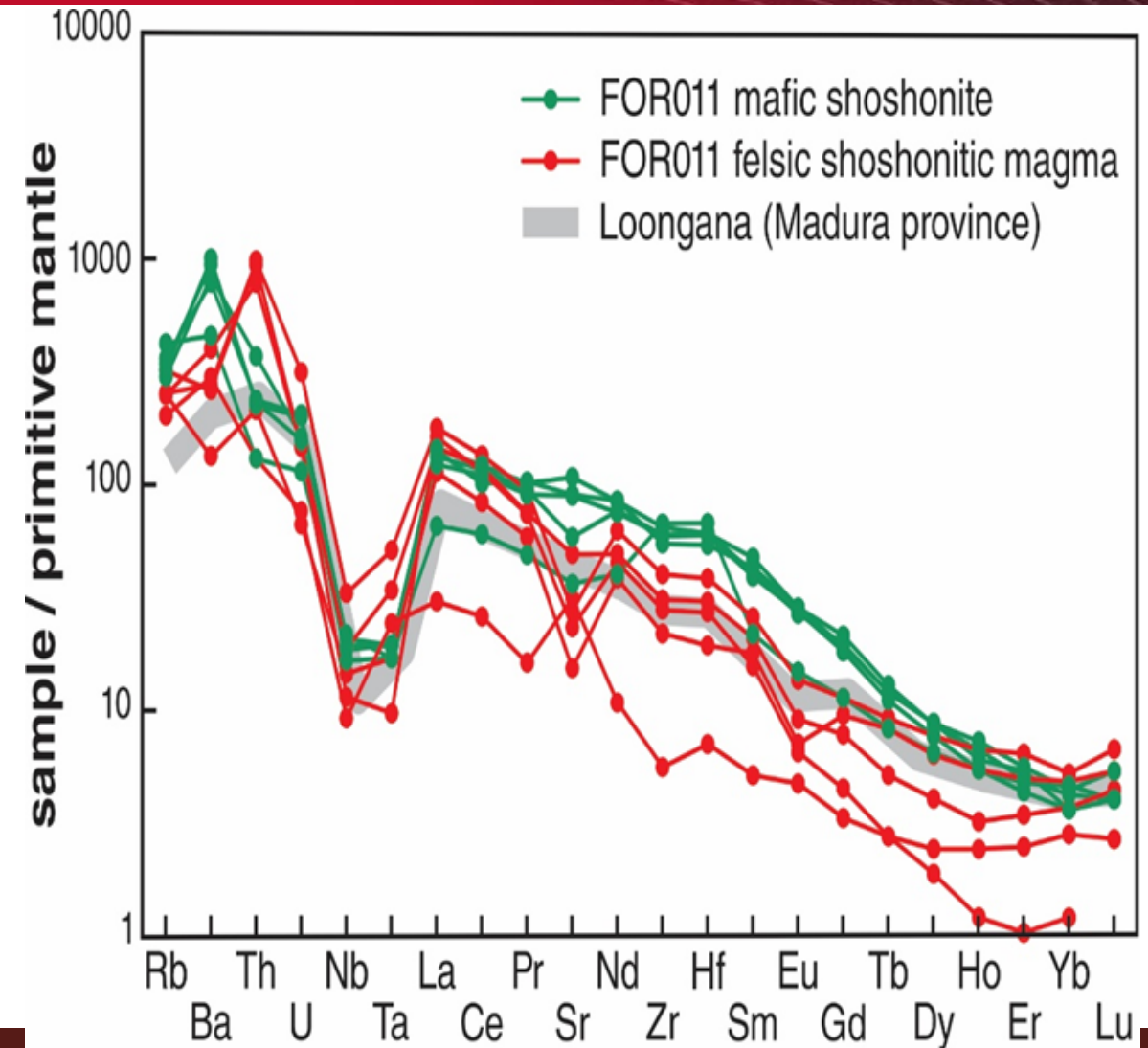
- Form a major component (up to 25%).
- Mafic (shoshonite) (SiO_2 48.45 – 60.87 wt%) and high silica (shoshonite series) (SiO_2 67.78 – 75.08 wt%) groups - each type occurring in both cores. Most isotopically juvenile magmas at that age.
- The combination of extremely high $\text{Mg}^\#$ and strong enrichments in incompatible trace elements points clearly to a lithospheric mantle source.



High-KMg series – melting of subduction metasomatised lithosphere



- This metasomatised lithospheric source was thick (within the garnet stability field) and wet and/or oxidized, and compositional evolution of the magmas was controlled by hornblende fractionation – high f_{H_2O} . These rocks show close compositional similarities with Cu-Au mineralised shoshonites in the Early Silurian Macquarie Arc (NSW).
- Single occurrence in the Madura Province represents melting at lower pressures (thinner lithosphere?)



Originally the same piece of mafic (oceanic) crust – related to the same ocean-opening event recorded in the AFO ocean-continent transition (relict E-MORB)

Pre c. 1200 Ma history of ocean closure/consumption and formation of mainly oceanic arcs – extensively in the Forrest Zone from c. 1720 to 1600 Ma; perhaps less extensively in the Madura Province – only recognised at c. 1400 Ma (although zircon inheritance also suggests an event at c. 1470 Ma)

Post c. 1200 Ma history of crust/lithosphere reworking (Great meltdown). Concurrent production of extremely contrasting magma types

- **high-KMg magmas from thick, wet, and oxidized subduction modified lithosphere mainly in the Forrest Zone;**
- **very high-T, dry and reduced high-KFe magmas in zones of highly thinned crust mainly in the Madura Province.**

As far as we can tell, high-KMg and -KFe magmatism is mutually exclusive in any one area. As with its occurrence in the Forrest Zone, the single sample of high-KMg shoshonite in the Madura Province is in 'arc-crust' at Loongana. High-KMg restricted to areas where thick lithosphere was developed (or remained!). Perhaps high-KFe magmatism indicates areas where thick lithosphere never developed, or where it was removed. Aeromagnetism should allow us to map these regions.